New geophysical and geological mapping of the eastern Baffin Bay region, offshore West Greenland

Ulrik Gregersen, Paul C. Knutz and John R. Hopper

The Geological Survey of Denmark and Greenland has carried out a detailed mapping project in the eastern Baffin Bay region covering c. 200 000 km² (Fig. 1). The purpose of the study was to update the previous mapping by using the most recent data and provide an improved basis for evaluation of the geological development and hydrocarbon potential of the region. After licensing rounds in 2007–2008 and 2010 considerable new geophysical and geological data were acquired in the region, and the extensive database now includes more than 100 000 km 2D seismic data and a number of wells (Fig. 1). The results of the work are summarised below and suggest deep basins and large ridges with complex structures.

Geological setting

In areas along the West Greenland continental margin, a number of basins with Proterozoic, Cretaceous and Cenozoic sedimentary successions have been identified (e.g., Dawes 1997; Dam *et al.* 2009; Rolle 1985). A number of rifted basins with large structural highs are interpreted to have developed in the east Baffin Bay region (Whittaker *et al.* 1997; Gregersen *et al.* 2013). During the Paleocene and Eocene, oceanic crust developed in central Baffin Bay and the Cretaceous rifted continental margin of West Greenland was separated from eastern Canada (Oakey & Chalmers 2012). The large-scale movements between Greenland and Canada generated new structures during the Palaeogene and reactivated faults within Cretaceous basins.

Results

The study included seven sub-projects: (1) seismic interpretation and mapping; (2) well correlation; (3) biostratigraphy; (4) seismic facies analysis and seismic inversion; (5) overview of source rocks; (6) maturity modelling; and (7) structural development. This paper describes some key results from the seismic interpretation and structural development. Interpretation of horizons and units were carried out with Schlumberger Petrel® software and included data from wells, seismic surveys, gravity surveys and magnetic

surveys. Additional data from seabed sampling and outcrops in the region were used for geological interpretation and as analogues. A robust framework was established with fourteen seismic stratigraphic horizons (A1–Hx) and eleven mega-units (A–H). Seismic cross-sections show deep, rifted basins separated by large structural highs (Fig. 2). The lateral extends and topography of the basins and structures are outlined in depth-structure maps (Fig. 3).

In the Kap York Basin (Fig. 4), the upper part of megaunit H can be correlated to parts of a 4.6 km/sec. TWT velocity zone from refraction seismic lines of Reid & Jackson (1997). They interpreted this zone to include Thule Supergroup sedimentary rocks, which crop out north of the study area (Dawes 1997). Thus the upper part of mega-unit H below horizon H1 (Fig. 2) probably includes sedimentary rocks with some analogues to the Thule Supergroup. Seismic interpretation suggests that some of the

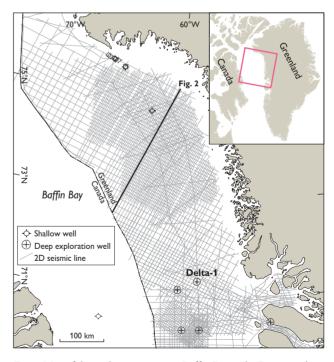


Fig. 1. Map of the study area in eastern Baffin Bay with 2D seismic data used and the location of shallow and deep exploration wells. Location of the seismic line of Fig. 2 is also shown.

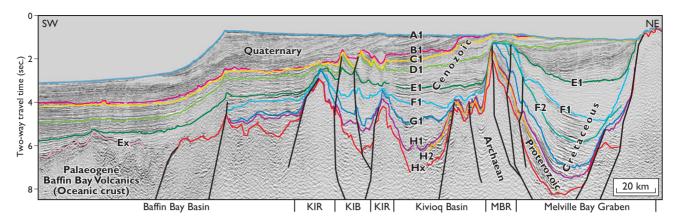


Fig. 2. Seismic line NE–SW across the eastern Baffin Bay region showing deep basins and large structures. The horizon boundaries of the seismic mega-units A–H are shown by coloured horizons from A1 to Hx and likely ages are shown. Note that some of the major deep-seated faults continue nearly vertically upwards through parts of the Cenozoic section indicating late compressional faulting. The names of the main structural elements are shown below the line and are from SW to NE: Baffin Bay Basin. **KIR**: Kivioq Ridge. **KIB**: Kivioq Basin. **MBR**: Melville Bay Ridge. Melville Bay Graben (Fig. 4). The seismic line is courtesy of TGS-NOPEC Geophysical Company ASA. Location of the seismic line is shown in Fig. 1.

wedge-shaped units in mega-unit H (Fig. 2) were probably formed by extension during the Proterozoic.

Sedimentary successions in the Baffin Bay may also resemble those from the Nuussuaq Basin, where Cretaceous nearshore to deltaic sandy deposits were documented (Dam *et al.* 2009). Seismic interpretation below horizon F1 (mega-unit F) shows prograding-aggrading clinoforms and troughs or channels possibly from deltaic systems (Gregersen *et al.* 2013).

During parts of the Late Cretaceous, relative tectonic quiescence prevailed and thick uniform units were deposited in the basins. Marine mudstone deposits with suggested source-rock intervals and oil seeps were recovered from both West Greenland (Bojesen-Koefoed *et al.* 1999) and northern East Canadian islands (MacLean & Williams 1983; Brent *et al.* 2013). Within the mapped basins, these marine mudstones may be analogous to deposits in mega-units E and F, between horizons E1 and G1 (Fig. 2). Extensional faulting occurred during the Late Cretaceous to earliest Paleocene by local rifting in the Nuussuaq Basin (Dam *et al.* 2009) and locally in a few other places in the region.

In north-eastern Baffin Bay, major Cretaceous rift basins trend SE–NW and are located east of the Kivioq Ridge (Fig. 4). West of the Kivioq Ridge, extensive volcanic areas have been mapped primarily from seismic reflection and magnetic anomaly data (Fig. 4). Studies with refraction seismic data (Damm 2010; Suckro *et al.* 2012; Altenbernd *et al.* 2014) showed oceanic crust in the eastern part of the Baffin Bay Basin. Seismic facies analyses combined with interpretation of magnetic data in the present study and in Gregersen *et al.* (2013) outline the eastern boundary of the

oceanic crust (also named Baffin Bay Volcanics) at nearly the same location (Figs 2, 4). The oceanic crust developed as Greenland and its rifted margin separated from Canada during the Paleocene and Eocene (Oakey & Chalmers 2012). The north-east and northward movements of Greenland's rifted basins caused compression-related

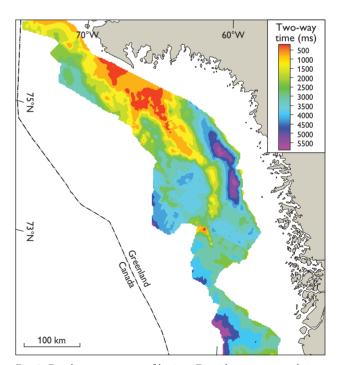
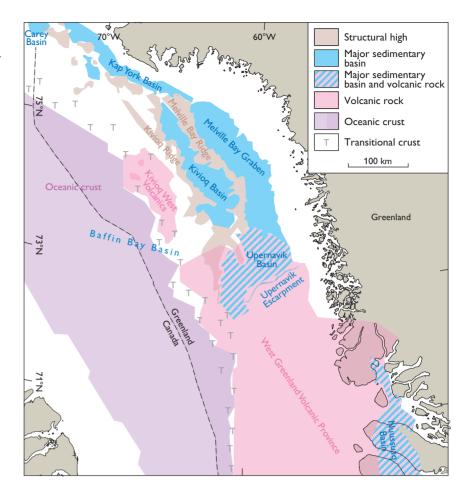


Fig. 3. Depth-structure map of horizon F1 with two-way travel time. This level is not drilled in the region but interpretation suggests that the horizon may be of a Late Cretaceous (Campanian?) age. The map illustrates the outline of large structures and basins of eastern Baffin Bay.

Fig. 4. Structural elements map of eastern Baffin Bay. The extent of the Canadian part of the Baffin Bay oceanic crust is from Oakey & Chalmers (2012).



tectonism with inversion, transtension and thrust faults developing (Fig. 2). The Palaeogene compression-related faulting in eastern Baffin Bay was probably a consequence of the same overall plate movements which also caused the Eurekan tectonic phase in North Greenland and Ellesmere Island. Portions of the large ridge structures likely contain block-faulted and post-rift Cretaceous sedimentary rocks later compressed into ridges such as the Melville Bay Ridge and the Kivioq Ridge (Figs 2-4). During the Palaeogene, flood basalts and other volcanic rocks developed in the West Greenland Volcanic Province and probably also in the Kivioq West volcanic area (Fig. 4). Paleocene-Eocene volcanic rocks cover parts of the Cretaceous basins such as in the Nuussuaq Basin (Dam et al. 2009; Larsen et al. 2015). Samples of flood basalt from the Delta-1 well offshore central West Greenland (Fig. 1) were dated to c. 50-56 Ma (Nelson et al. 2015). Sedimentary geometries suggesting basin-floor fan deposits were formed during the late Eocene to Oligocene in mega-unit D, above horizon E1 (Fig. 2), presumably related to the Eurekan compressive phase. This was followed by more passive infilling of the remnant rift basins lasting until middle Miocene (horizon

D1; Fig. 2). During the late Miocene to Pliocene, large contourite drift- and mass-transport systems (e.g. mega-slides) developed in Baffin Bay within mega-units B and C (between horizons B1 and D1; Fig. 2) (Knutz *et al.* 2015). The final phase of basin development is characterised by major progradation of the West Greenland shelf towards southwest (mega-unit A between A1 and B1; Fig. 2) mainly as a result of trough-mouth fans formed by late Pliocene–Pleistocene glaciations (Knutz *et al.* 2015).

Interpretation of seismic facies and attributes has identified a number of amplitude anomalies (bright events) and vertical disturbances (chimney structures) on seismic data over structural highs in the Cenozoic section. Some of the features could relate to upwards migration of fluids and possible active petroleum systems. Some of the source rocks that produce oil seeps in the Nuussuaq Basin and equivalent sources in west Baffin Bay could occur in basins of eastern Baffin Bay (Bojesen-Koefoed 2011). Large structural crests within Cretaceous sections occur close to deep basins (Fig. 2), and some of the structures may include potential traps for hydrocarbons, given the right conditions.

The updated mapping and new interterpretation in this study have improved the large-scale mapping of the structures and basins of the eastern Baffin Bay region, outlining a prospectivity potential and also pointing to uncertainties and risks that require future clarification.

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